GROUND WATER DEPTH PREDICTION FROM SATELITE DATA

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in Partial Fulfilment of the Requirements
for the Degree of
MASTER OF TECHNOLOGY

By
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INDIAN INSTITUTE OF TECHNOLOGY, KANPUR

DECEMBER, 1980



- 1 SED 1981

To all those who made the 'ARYABHATTA' and the 'BHACKARA' possible

CERTIFICATE

This is to certify that the work GROUND WATER DEPTH PREDICTION FROM SATELLITE has been carried out under my supervision and has not been submitted elsewhere for a degree.

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SYNOPSIS

Dissertation on

GROUND WATER DEPTH PREDICTION

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This project explores the possibility of shallow ground water depth prediction from satellite data. Two methods are used for this purpose the computer compatible tape method and the Micro-densitometer method.

Both methods are based on the fact that reflectance characteristics of a particular soil cover vary with variations in the ground water depth. Calibration curves between reflectance units (CCT method) or light density (densitometer method) and water depth, are developed from the field it and these curves are used for prediction of ground water depths. The CCT method is more accurate than the other method. It was found that these methods are not applicable for water depths over 10 m, or for hilly or built up regions. Very encouraging results were obtained in

The prediction of ground water depths in alluvial soil (Gujrat region) in depths less than 10 m. For such regions this study establishes that the for making ground water contour maps of large areas, the most economical and convenient method is a satellite survey.

INTRODUCTION:

1.1 <u>Basic Principles</u>: The objective of this thesis is to explore the possibility of predicting shallow water depths from satellite data. This study is aimed at providing a more economical and convenient medium (i.e., a satellite survey) compared to ground surveys for making ground water maps of large areas.

The area selected for study encloses a part of Rajasthan and a part of Gujrat (about 185 Km. x 185 Km.). A broad geological classification of this area is given in Fig. 1. The satellite data pertaining to this specific area was obtained by LANDSAT in the form of a black and white photo negative, and Computer Compatible Tapes (CCT). This satellite data - the imagery and two CCTs - were procured by the Indian Space Research Organization (ISRO), from the EROS Space Centre, South Dakota, USA.

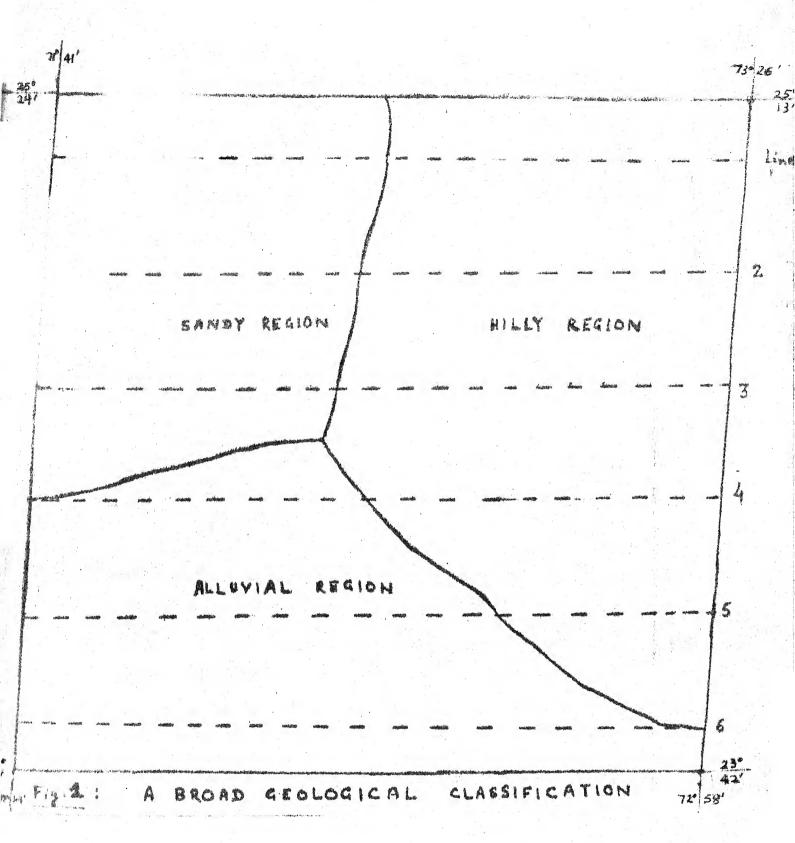
Remote sensing is the science of obtaining khowledge without physical contact. The remote sensing satellite by the use of Multi-Spectral Systems (MSS) can gather such 'knowledge' from an altitude of several hundred kilometres (920 Km. for LANDSAT), by scanning the land masses underneath and recording the neglected solar radiation (or 'reflectance') in different wave-bands. For water studies, we use the reflectances recorded by the MSS in the infra- red region or band - 7 (0.8-1.1 m wave- length) because in this 'and water appears black on the positive print due to strong a sorption bands in the near Infra-red part of the spectrum. For a particular soil in a given area, the variations in shallow water

depths alter the tone of the ground surface and these variations, most discernible in Band - 7, alter the reflectances of the same ground surface depending upon the proximity of the ground water to the ground surface. Hence the variation in reflectances, of the same ground surface is related to the variation in the ground water depths. By measuring the ground water depths by bore holes etc. at a number of distributed points in the given area and by recording the corresponding reflectances from the available satellite data, we can study the correlation and plot a curve of reflectance versus ground water depth. This is a calibration curve, and it can also be called a calibration - prediction curve because it can be used to predict the ground water depths at different points in the given area knowing the corresponding reflectances.

Two techniques based on this principle have been employed in this research study for the prediction of shallow ground water depths. One uses a micro-densitometer and the other is based on the application of magnetic tapes (CCTs). The CCT approach required an extensive use of the DEC-10 computer available in IIT Kanpur.

The promising results obtained in this study highlights the scope of satellite application for predicting shallow water depths. For a large area survey like a state survey or a national survey for marking ground water depth contours, it is obvious that the most economical and convenient option is a satellite survey.

* * *



MAGNETIC TAPE METHOD:

2.1 Computer Compatible Tape (CCT) Approach:-

The data collected from a Multi spectral scanner (MSS) on a aircraft or satellite can be stored on a Computer Compatible

Tape after being processed at a ground data collection centre. The

MSS system collects ground reflectance data by scanning along lines

perpendicular to its flight path and recording reflectance values

as 'pixels' along each 'line'. The resolution of a pixel in Landsat-3

is about 80 m. After being processed at the ground station to

eliminate undesirable interferences or 'noise', the data or 'signals'

is recorded on a CCT in the form of 'records', The 'pixels' or reflectances

are recorded in 'bits' on a CCT in the binary mode.

2.2 CCT Storage Mode:

Information is stored in a CCT in the form of 'records'. Each record represents one line scanned by a satellite in a given imagery. Each record is subdivided into 'words' - a smaller storage unit, i.e. And each word is subdivided into 36 'bits', a bit being the smallest storage space on a CCT. Each bit stores one binary number. In our case, one word in the CCT Stores four pixels, each pixel occupying seven bits of space (hence the maximum reflectance value cannot exceed 2 = 128). The reflectance values range from 0 to 128.

As the MSS system records data in Bands 4, 5, 6 and 7. The same pixel is recorded on a '800' density or '1600' BPI tapes in four different parts in the four wave ands. Each '800' density

CCT contains two sections or 'files', and each file contains data of the same one imagery in a given wave band. Thus to record the same imagery in Bands 4, 5, 6, 7 seperately, two '800' density CCTs are required, each wave band confined to one file on a tape.

The CCT containing Band - 7 part of the imagery was used for ground water study in this study because variations in ground water profile are most discernible in Band - 7 (0.8-1.1 pm wavelength):

A Computer program (see Appendix No. I) was developed which gives the line number and the pixel number of a particular spot on the imagery or a band map of the given Rajasthan - Gujrat region for the given latitude and longitude coordinates of this particular spot.

Hence, to read the reflectance value from a CCT at any given point on a imagery, first the latitude and longitude of this point are determined from the topo sheet and these coordinates are fed into the computer program to get the required line number and pixel number of the given point. Once the line number and pixel number are known, the CCT is used to read the records on the tape till we reach the required record (or line scanned by the MSS)

Next we activate computer program (see Appendix II) which reads the record, converts the data from binary to decimal mode and prints the given reflectance (or 'pixel' scanned by the MSS).

2.3 Central pixel Method & Grid Averaging Method:

Four sets of correlation were attempted between reflectance the so and ground water depth. In the first case or / called the 'central pixel method', only the exact pixel corresponding to the given latitude and longitude coordinates of a given point on a land map or imagery,

was selected and its reflectance value was correlated with the corresponding water depth.

In the Grid Averaging Method, we take a number of pixels on either side of the central pixel and take the arithmatic mean of the reflectances of these pixels which form a square or grid with the central pixel occupying the centre position. It is obvious that for a square grid, each side/the grid must have an odd number of pixels. Hence we can have a grid with 3 pixels on each side, 5 pixels on each side, 7 pixels on each side etc. In this study these three grid values were chosen (3, 5, 7 pixels, i.e.). "t can easily be seen that in the case of the grid approximation method for 3, 5, 7 pixels making up the grid sides we will have to take the arithmetic mean of 9, 25 and 49 reflectance valves respectively to get the representative reflectance of each grid. For a large number of control pixel points, manual computations being time consum ing for grid averaging at each point, a simple computer program was used which gave the representative reflectances for both central pixel method and Grid Averaging Method (5, 6, 7 pixels grids) for each observation point.

After this step is completed, knowing the latitude and longitude of a given observation point, four representative reflectances corresponding to this observation point (one for the central pixel, the other three for the 5, 6 and 7 pixels size grids) are computed. The next step is the preparation of reflectance versus depth curves.

The four representative reflectances, obtained from the previous step, were plotted individually against observed ground water

correlations were attempted. After the points were marked on each graph the next step was curve fitting. A qualitative inference about the nature of the calibration curve for the magnetic tape approach, was deduced from basic principles. As water appears black in Band 7 the shallower the water table is the more affected the ground surface would be from the moisture proximity.

Hence for lesser water depth the same ground surface will give lesser reflectance and for deeper depth the reflectance will be greater. For a surface water body like a lake, the reflectance will be zero, the minimum reflectance value. Hence we have an exponential growth curve of increasing slope, between reflectance values and the corresponding water depths. Mathematically, the equation of such a calibration curve fitted is:-

$$y = e^{bx+c}$$
 - (II)
where 'b' and 'c' are constants,
 $y = reflectance$,
 $x = ground water depth$

In order to fit a calibration curve obeying Equation (II) in the four sets of graphs between the representative reflectances and the observed ground water depths, the method of least squares was adopted.

Evaluation of 'b' and 'c' in Equation (II):

Taking the natural logarithm of both sides in Eqation (II), $log_e y = bx+c$

Let
$$Y = log_e y$$
, $b = BB$, $c = cc$
 $Y = BB x + CC$

It can easily be established that the normal equations are,

$$\mathbb{Z}Y = \mathbb{B}\mathbb{B}\mathbb{Z}\mathbf{x} + n \ \mathbb{C}\mathbb{C}$$
 - - (III)

$$\mathbf{x}\mathbf{x}\mathbf{y} = \mathbf{B}\mathbf{B}\mathbf{x}^2 + \mathbf{C}\mathbf{C}\mathbf{x} - - - (\mathbf{I}\mathbf{v})$$

where n = number of observed points.

The values of 'BB' and 'CC' can be found by simultaneously solving Equation (III) and (IV) by inserting the observed values of 'X' and 'Y'.

Once 'BB' and 'CC' are known the standard error, , for the "fit" determined from:-

$$\sum_{i=1}^{2} (x_i - \overline{x})^2$$

After the 3-6 test had been applied, the coefficient of correlation was determined to find whether there was any correlation between 'x' and 'y' and if so, how good it was. The correlation coefficient is defined as: $\sum (x-\bar{x}) \ (y-\bar{y})$

 $r = \frac{\sum (x - \bar{x}) (y - \bar{y})}{\sqrt{\sum (x - \bar{x})^2 \sum (y - \bar{y})^2}}$

The value of 'r' can vary from -1 to +1. A negative 'r' indicates that 'y' decreases with increase in 'x'. If r=0, it indicates an absence of correlation. If r=0.5, a weak correlation exists; if r=0.8-1.0, a good correlation is concluded (subject to the usual standard statistical tests with the given level of confidence).

Results: In this study, with the help of a computer program (see AppendixIII), the calibration curve equations, standard deviations and coefficients of correlation were determined for the four plots in the magnetic tape method for the Gujrat region. These results are given in Table - 1.

Table - 1

METHOD	CALIBRATION CU	JRVE EQUATION BBx+cc	STANDAR DEVIATI		CORRELATION COEFFICIENT
4. 1.	ВВ	CC	for'x'	for'	r' 'r'
Central Pixel 3 pixels grid	0.20175720 0.16010590	1.92411720 2.41294530	2.047 2.047	0.4	40 0.9389 680.8916
5 pixels grid 7 pixels grid	0.07443128 0.03375640	3.08453320 3.46898590	2.047	0.17	
i bixono gi id		1	1	!	

The field data for the Gujrat region used for obtaining these results are given in Table 2(Appendix IV).

The field data for the Rajasthan region - used in making reflectance vs. depth plots (see fig. 2-5~) by central pixel method and 3 pixels, and 5 pixels linear averaging methods - are given in Table 3.

No appreciable correlation was found between reflectance and depth in the Rajasthan region after inspecting figs. 2-5.

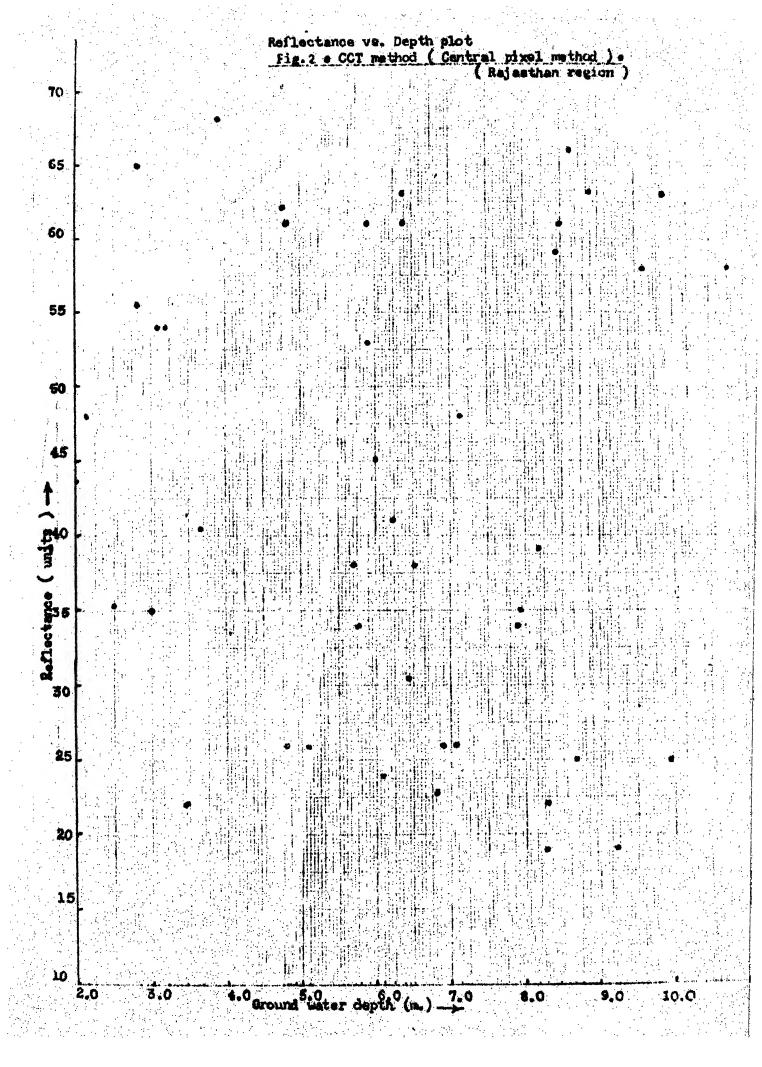
Sl. no. Longitude (Degrees)		Line No.	Pixel No.	Reflect Central pixel method		5-pix. grid	7-pixel grid	Depth (m.)
1, 72,850000	24.340000	1436	2195	19	27	31	35	5,0
2. 72.425000	23.915000	2163	1764	59	57	45	42	10,0
3. 72.366666	24.266666	1 656	1524	21	24	31	35	5,8
4. 72.366666	23,916666	2174	1679	34	43	42	48	8.7
5. 72.020000	24,520000	1357	910	27	26	32	39	6,8
6. 71.810780	25.354423	172	241	58	59 .	47	43	9.9
7. 71.866667	24,300000	1721	781	24	39	41	45	7.7
8, 71,625000	25,363000	437	167	19	22	30	36	4.8
9. 71. 6. 6.	24.067376	2048	1171	24	37	40	47	5,4

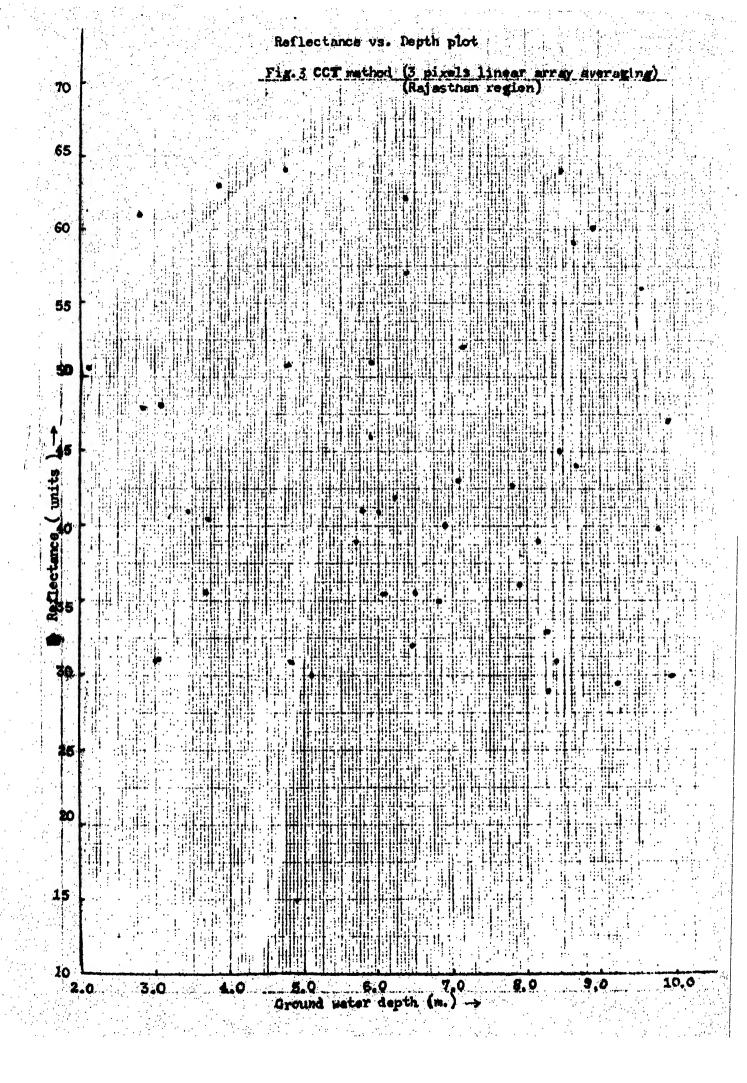
Table - 3 Field data - Rajasthan region (Longitudes and Latitudes of these field points are given in Appendix V.)

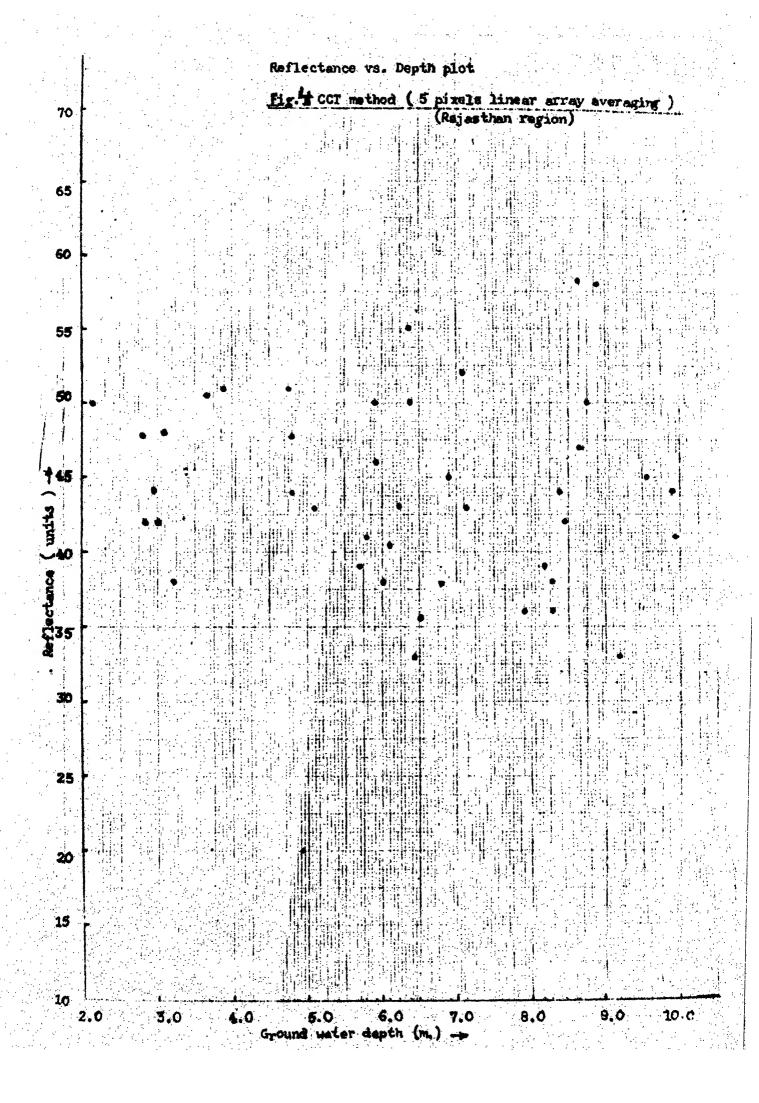
	·	In A		V 🕯)	en de la companya de La companya de la co		
Sl.	Line	Pixel	REFLE	CTANCE			
/ No.	No.	No.	[Central	3 pixels	5 pixels		Measured depth
	<u> </u>	Q	≬ pixel	l average	average	laverage	(m.)
			2				
1.	179	242	63	47	44	42	9,90
						177	
2,	235	9 28	22	41	44	47	3 <u>,</u> 46
2	226	4000		5 0	÷ o	4.0	0.64
3.₄	236	1328	66	59	58	49	8,64
4.	248	1256	4 İ	42	43	49	6.23
- •			•	7.43	10	22	0 & 20 0
5.	261	1179	25	44	47	48	8.66
		_					
6.	271	1296	58	51	50	42	2.11
-T	270	761	2.4	30	4 +	42	E E
7.	278	764	34	38	41	41	5,72
8.	285	1208	59	45	42	41	8,45
9.	286	825	27	40	45	47	6,90
10,	299	1288	39	39	39	44	8.16
11.	328	607	62	52	50	47	5,90
	7						
12.	337	609	35	31	42	46	2,99
13.	338	7 7 6	38	39	39	45	5,69
14.	340	644	48	52	43	39	7.72
15.	351	849	68	63	51	4 6	3,38
					·		
16.	385	935	90	29	36	39	8,27
17.	419	630	45	41	3 8	32	3,00
18,	426	1136	63	57	55	47	6,38
19.	436	883	22	33	38	42	8.27
20.	438	167	27	31	44	47	4.80
21.	447	185	61	62	50	45	6,39
	20.0	above and					

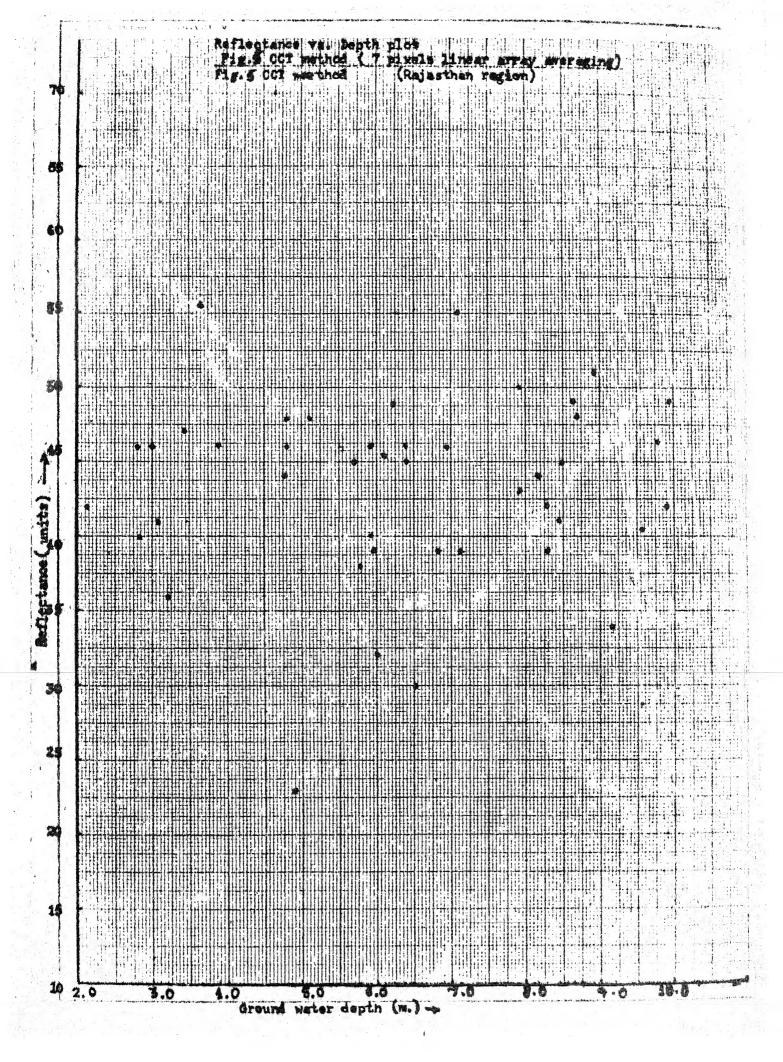
Contd. Table - 3

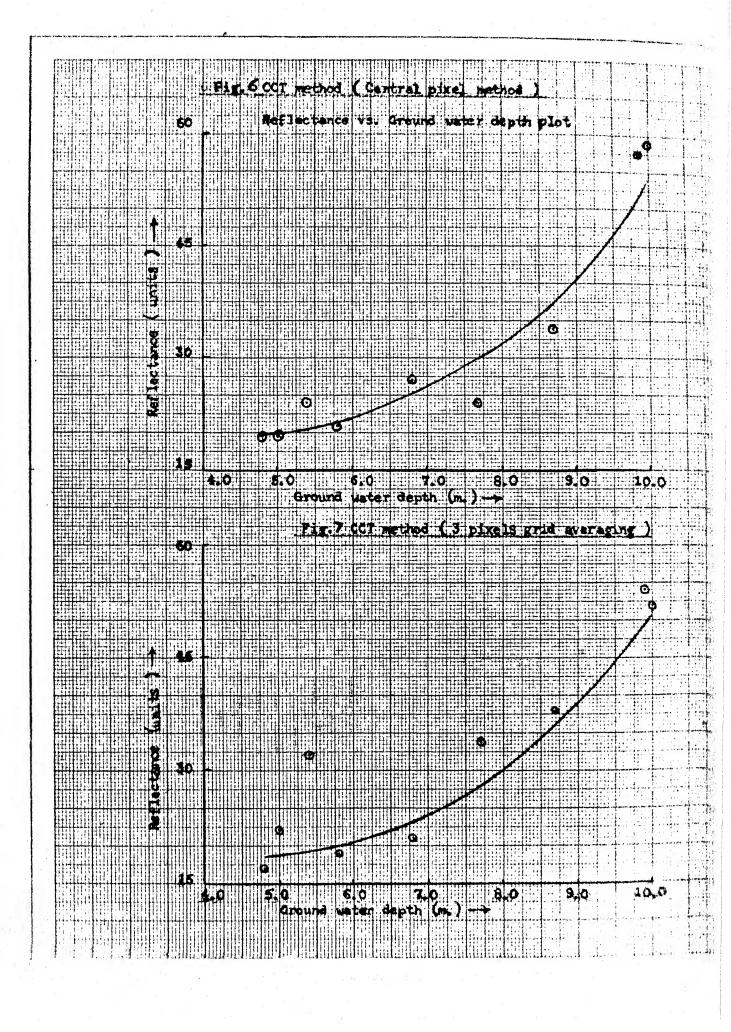
Sl.	Ž .	<u> </u>		FLECTANCE		Į	Measured
No.	[Line No.		[Central [pixel		5 pixels		depth
	Ĭ NO.	, 140.	Y DIXEL	laveravge L	I average	laverage l l l	(m,)
22.	455	1035	6.9	68	58	53	8,75
23.	456	754	56	46	42	40	2.77
24.	460	675	54	49	48	41	3,06
25.	467	1283	63	- 60 -	58	52	8-90
26.	483	909	26	43	52	55	7.80
27.	489	734.	. 65	61	49	46	2,80
28.	479	631	62	64	53	44	4.75
29.	5 09	1237	31	32	33	39	6.45
3,0.	5 1 9	1177	19	29	33	34	9.20
31 🌲	525	724	25	30	42	49	9.95
32.	533	1168	38	36	36	30	6,50
33,	553	737	41	36	51	56	3.67
34,	563	1232	53	47	47	40	5.89
35,	629	435	24	36	42	46	6.10
36,	646	692	27	30	43	48	5,10
37.	7 09	216	54	41	38	37	3,20
38.	900	414	58	57	45	42	9,53
\$ 9.	1251	314	9	15	. 20	23	4,90
40.	1360	9 09	23	35	38	39	6.80
41.	1682	402	62	53	48	49 ,	4.80
42.	2047	776	34	37	37	43	7.90
43.	2127	494	35	32	44	50	8.40

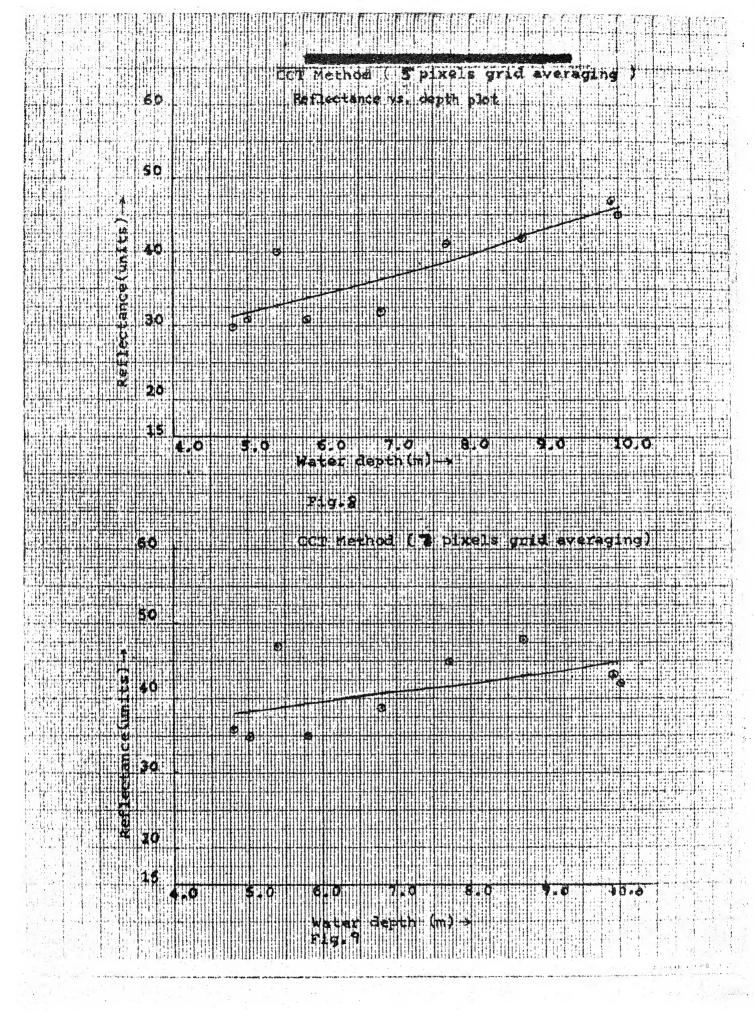












2.5 Prediction Process (Magnetic Tape Method):

Prediction Method: A number of points were taken on a 1" = 4 miles scale ground water contour map of the given region and their latitudes and longitudes were determined. Using these latitudes and longitudes in the 'LATPIK' computer program (see Appendix I), the line numbers and pixel number corresponding to each point were found. Then using these line numbers and pixel numbers and the 'TAPE. FOR' computer program (see Appendix II) the corresponding reflectances (pixel values) were found.

Next with the help of the calibration curves and these reflectances the corresponding ground water depths were predicted for the given set of points taken on the map. These predicted ground water depths were then compared with those indicated in the conventional ground water contour map to complete the prediction process.

To attempt a prediction of ground water depths in the Gujrat region (lower half of the given imagery) the Banaskantha district which mostly covers this part was selected. This district area was subdivided into three sub regions (see Fig. 15) for analysis. Important details about each region are given below:

- Region (1) Mostly alluvial with ground water depths 10 m

 Latitude

 Longitude 71.45 71.85, Latidue: 24.00 24.50
- Region (2) Alluvial transition to hilly region, water depths > 10 m Longitude : 71.90 - 72.30, Latitude : 24.05 - 24.50
- Region (3) Mostly hilly region with alluvial regions in minority.

 Depths: variable; includes > 10°, depths also.

 Longitude: 72.35 72.70, Latitude: 24.10 24.25

Region (1) was expected to be most suitable for prediction purposes and the results of the two prediction methods - tape method & densita meter method bore this out. In Region (2) & Region (3) the methods are

unreliable because of the deep depths and hilly terrain.

RESULTS AND DISCUSSION: The field data etc. isgiven in tables 4-7.

A separate case study was taken up to show that the satellite survey methods are not useful for detecting depths deeper than 10 m. Ninetten field measurements of water depths greater than 10 m were undertaken and the prediction process using the tape method was applied (See table - 7). The following results were obtained:

N = 19 points, O = 2.75 m

The 3 - σ test shows that we have to reject any x > 8.27 m hence we reject serial no. 16 observation table - 7. Therefore, we finally get:

N=18,

Standard error, 0 = 2.63 m

The high value of the standard error proved the point adequately that satellite survey methods are unreliable in depths greater than 10 m.

Region (1) : (See table - 4)

Number of observation points, N 61 Standard error, $\sigma_n = \frac{1}{2} \cdot 1.36 \text{ m}$

The 3 - σ test shows that we have to reject any x > 4.08 m hence we reject the following two observations: x = 7.84 m (Longitude: 71.6° , Latitude = 24.4°) and x = 6.13 m (Longitude = 71.75° . Latitude = 24.15°).

After removing these two points we get: N=59, $\sigma_n = \frac{+}{-}$ 0.99 m. The standard error is \$\frac{1}{2}\$ 0.99 m. The tape method thus gives a good result in Region (1), which is mostly alluvial in mature and with water depths less than 10 m, where our calibration - prediction equations yield the best results.

Region (2): (See table - 5)

N = 61 $\sigma = 4.23$ m

The 3 - 0 test shows that we have to reject any x > 12.69 m, hence we reject the ten observations corresponding to serial nos 28, 31, 36, 37, 38, 42, 43, 46, 48 and 52 in table - 5.

After removing these ten points we get :

N = 51,

Standard error, G = 3.39 m

Therefore, the standard error being 3.39 m, we can conclude that in Region (2) we get highly unreliable results. This region is mostly of more than 10 m ground water depth, where our satellite survey methods are not applicable.

Region (3) : (See table - 6)

N = 10, G = 1.97 m

The 3 - 0 test shows that we have to reject any x > 5.91 m, so we reject one observation corresponding to serial no. 10 in that table - 6.

Therefore, we have :

N = 9.

Standard error, 0 = 1.84 m

We observe that a few points (for ex., serial no. 3 in table - 6) show large differences between predicted and interpolated depths.

This occurs most probably because the nature of the local terrain, which is hilly (extension of Mount Abu hill region). A few points show a very good agreement between interpolated and predicted depths (for ex., serial no. 9) and this could be attributed to alluvial patches in between the part of the hilly regions for ground water depths less than 10 m.

 $\frac{\text{Table}-4}{\text{Field data \& calculations for prediction}}$ (CCT method: Region (1), Gujrat)

Sl.	Longitude (Degrees)	 Latitude (Degrees)	Inter- polated Depth,D (m.)	ance	Predicted depth, P (m.)	<pre>Deviation x = {D-P; (m.)</pre>
1.	71.45	24,20	5.00	27	6,80	1.80
2.	71.45	24.05	9.00	23	6.42	2.58
3.	71.45	24.00	9:00	59	10.67	1.67
4.	71,50	24.25	5:00	22	5.78	0.78
5.	71.50	24.20	6 ; 48	33	7.79	1,31
6.	71.50	24.10	6.81	36	8.23	1.42
7.	71.55	24.30	4.00	23	6.00	2.00
8.	71.55	24.20	6.67	58	10.59	3 . 91
9.	71.55	24.15	7.00	58	10.59	3.59
10.	71.55	24.10	7.00	51	9.95	2.95
11.	71.55	24.05	8.59	61	10.84	2.25
12.	71.55	24.00	8.18	32	7.64	0,54
13.	71.60	24.40	3.00	61	10.84	7.84
14.	71.60	24.35	4.33	24	6.22	1.84
15.	71.60	24.30	7.00	28	6.80	0.20
16.	71.60	24.25	9.00	72	11.66	2.66
17.	71,60	24.20	7.00	24	6.22	0.88
18.	71.60	24.15	7.00	26	6,61	0.39
19.	71.60	24.05	7.48	38	8.49	1.01
20.	71.65	24.05	7.06	26	6.61	C _• 45
21.	71.65	24.45	5,64	45	9,33	3,69
22.	71.65	24,40	4:80	25	6.42	1,62
23.	71.65	24.25	7.00	22	5 . 78	1.22
24.	71.65	24.20	7.00	48	9.65	2,65
25.	71.65	24.15	7.00	33	7.77	0.77
26.	71.65	24.10	7.00	38	8.49	1,49
27.	71,65	24,00	5.83	20	5 .7 8	0.05

The second second second						*
Sl. No.		Latitude (Degrees)	Inter- polated Depth,D (m.)	 Reflect- ance	Predicted depth, P (m.)	Deviation (x'= D-P (m.)
28.	71.70	24.50	7.44	27	6.68	0.74
29.	71.70	24,45	7.29	20	5.31	1,98
30.	71.70	24.40	6.47	25	6,42	0.05
31.	71.70	24.25	7.00	52	1.0.05	3,05
32.	71.70	24.20	6,08	20	5.31	0.77
33.	71.75	24.50	8.02	21	5,55	2.47
34.	71.75	24.45	. 7.77	29	7.15	0,62
35。	71.75	24.40	. 7.51	52	10.05	2.54
36.	71.75	24.35	7.24	17	4.51	2.73
37.	71.75	24.30	7.00	45	9.33	2,33
30.	71.75	24.25	7.00	27	6.80	c. 20
39.	71.75	24.20	5.00	26	6.61	1,61
40.	71.75	24.15	₄ 5.17	67	11.30	6.13
41.	71.75	24.10	5.04	25	6.42	1.34
42.	71.75	24.00	5.00	29	7.1 5	2,15
43.	71.80	24.50	8.65	25	6.42	2.25
44.	71,80	24,45	8.48	44	9,22	0.74
45.	71.80	24,40	8.28	32	7.64	0,64
46.	71.80	24.35	8.00	, 25	6,42	1.58
47.	71,80	24.30	9. 25	37	8,36	0.89
48.	71.80	24.25	9.63	38	8,49	0.14
49.	71.80	24.20	9,00	49	9.75	0.75
50,	71.80	24.15	7.80	40	8.75	0.95
51.	71.80	24,05	7.48	38	8,49	1.01
52.	71.80	24.00	5.10	33	7.79	2,68
53,	71.85	24.05	9,24	50	9.85	0.61
54.	71.85	24.45	9 .2 8	26	6.61	2.67
55.	71.85	24.40	9,16	28	6.98	2.18

Contd. Table - 4

Sl. No.	Longitude (Degrees)	Latitude (Degrees)	Inter - polated Depth,D (m.)	Reflect- ance	Predicted Idepth, P I (m.) I	Deviation x = \D-P (m.)
ĪĠ.	71.85	24.35	9.33	50	4.85	0,52
57.	71.85	24.30	11.57	38	8.49	3,08
58.	71.85	24.25	12.20	65	11.15	1.05
50.	71.85	24.15	12.50	. 70	11,52	0.98
,60.	71.85	24.10	13.00	61	10.84	2.16
61.	71 _• 85	24.05	7.00	35	3, 9 8,09	1.09

Table - 5

Field data & calculations for prediction (CCT method : Region (2), Gujrat)

Sl. No.	Longitude (Degrees)		Inter - polated Depth,D	Reflect- ance	Predicted depth, P (m.)	Deviation X = ;D-P; X (m.)
.1.	71.90	24.50	9,90	24	6,22	3,68
2.	71.90	24.45	10.14	23	6.00	4.14
3.	71.90	24.40	10,58	53	10.14	0,44
4.	71.90	24.35	11.77	32	7. 64	4.13
5 👞	71.90	24.30	13.75	38	8.49	5.26
6.	71.90	24.25	14.76	25	7.42	8.34
7.	71.90	24.20	14,76	62	10.92	3.84
8,•	71.90	24.15	14.80	76	11.93	2.87
9.	71.90	24.10	15.0	32	7.64	7.36
10.	71.90	24.05	11,40	61	10.84	0.56
11.	71.90	24.00	7.50	38	8.49	0.99
12.	71.95	24.50	10.53	29	7.15	3.38

Contd. Table - 5

Sl.	Longitude (Degrees)		Inter - polated Depth, D (m;)	Reflect- ance	Predicted depth, P (m.)	Deviation $x = D - D$ (m_{\bullet})
13.	71.95	24,45	10,75	47	9,55	1.20
14.	71,95	24.40	12.33	31	7.48	4.85
15.	71,95	24.35	12.88	38	8.49	3.59
16.	71.95	24.30	14,27	18	4.79	9,48
17.	71,95	24.25	19,33	65	11.15	6.18
18.	71,95	24.20	17.33	22	5 .7 8	9.10
19.	71,95	24,15	17.27	36	8.23	9.04
20.	71.95	24.10	14.50	59	10,67	3.85
21.	71.95	24 05	10.25	27	6.80	3.45
22.	71,95	24.00	10.20	62	10.92	0.72
23.	72,00	24.50	15.00	47	9.55	5.45
24.	72.00	24.45	13.00	26	6.61	6.39
25.	72.00	24.40	12.50	42	8.99	3.51
26.	72.00	24.35	14.20	17	4.51	9.69
27.	72.00	24.30	18.57	27	6.80	11.87
28.	72.00	24.25	19.70	14	16.16	12.83
29.	72.00	24.20	19.66	36	8.23	11.43
3⊅.	72.00	24.15	17.67	64	11.10	6.56
31.	72.00	24.10	14.40	8	0.77	13,63
3₽.	72.00	24.05	10,50	45	9,33	1.17
33.	72.05	24.50	19.00	29	7.15	11.85
34.	72.05	24.45	19.00	27	6.,80	12,30
35.	72.05	24.35	21.00	46	9.44	10,66
36.	72.05	24.30	22.00	26	6,61	15.39
37.	72.05	24.25	22.00	26	6.61	15.39
33.	72.05	24.01	17.00	13	3.18	13.82
39.	72.05	24.10	14.27	68	11.30	2.89
40.	72.10	24.40	19.00	* 26	6,61	12,39
41.	72.10	24,35	19.33	56	10.42	8.91
42.	72.10	24.30	20.67	27	6.80	13,97
43.	72.10	24.25	22.61	36	8.23	14,39

Sl. No.	 Longitude. (Degrees) 	(Latitude (Degrees) 	Inter - Ipolated IDepth, D I (m.)	Reflect - ance		<pre>Deviation x = D -P (m.)</pre>
44.	72.10	24.15	16.78	76	11.93	4.86
45,	72.10	24.10	14.67	26	6,61	8.06
46.	72.15	24,35	19.00	23	6,00	13,00
47.	72.15	24.30	19,40	49	9.75	9.65
48.	72.15	24.25	19.90	22	5.78	14.12
49.	72.15	24.20	18.75	26	6.61	12.14
50.	72.15	24.15	16.05	50	.9.85	6.19
51.	72.15	24.10	14.53	36	8.23	6.31
52.	72.20	24.30	19.00	19	5.06	13.94
53.	72.20	24.25	19.00	53	10.14	8.86
54.	72.20	24.00	18.433	66	11,23	7.10
55.	72.20	24.15	16.14	35	8.09	6,29
56,	72.20	24.10	14.11	66	11,23	2.88
57,	72.25	24.25	17.50	53	10.14	7.36
58.	72.25	24.20	18.23	34	7.94	10.29
59.	72,25	24.15	15.95	47	8,36	7.59
60.	72,30	24.25	13,00	23	6.00	7.00
61.	72.30	24.20	16.11	59	10,67	5.44

Table - 6

Field data & calculations for prediction (CCT method : Region (3), Gujrat)

-		Χ				
Sl.	Longitude (Degrees)	Latitude (Degrees)	Inter - polated Depth, D	l Reflect - l ance l	Predicted Idepth, P (m.)	Deviation x = D -P (m.)
				<u> </u>		
1.	72.35	24,25	7.57	44	9.22	1.65
2.	72.35	24.20	±0.17	20	5.31	4.86
3.	72.40	24.25	5.00	59	10.67	5.67
4.	72.40	24.20	500	29	7.15	2.15
5.	72.45	24.15	9.80	26	6,61	3 .1 9
6.	72.50	24.15	9,00	50	9.85	0.85
7.	72.55	24.15	11.00	29	7.15	3,85
8.	72.60	24.15	11.00	. 28	6.98	4,20
9.	72.65	24.15	11.00	59	10.67	0.33
10.	72.70	24.15	11.00	19	5,06	5,94

Table - 7

Field data & calculations (CCT method :Gujrat region; for depths > 10 m.)

							300 0 1 V	and the same of
SI.	Latitude (Degrees)	Longitude M	easured depth (m.)D	Line No.	Pixel No.	Reflect ance	Predicted Depth(m)P	Deviation x= P-D; (m.)
1.	24,280: .,	72 , 180′)	11.50	1679	1247	51	9,95	1 •55
2.	24,220	72,020	12.20	1804	1040	48	9,65	2.55
3.	24.350	72.530 ^0	10.80	1495	1725	51	9,95	0.85
4	24.190	72,430′	11.20	1755	1650	59	10.59	0,61
5 🛊	23.875 .	72.635: 3	12.00	2173	2089	48	9,65	2.35
6.	23,850	72,640'	12.00	2209	2107	27	6.70	5.30
7.	23.825	72.645	12.00	2245	2125	17	4.51	5.49
8.	23,850	72,665	14.00	2203	2143	36	8.23	5.77
9.	23,825	72,675	14.00	2238	2169	24	6.24	7,78
10.	23 ₄ 800'	72.680.	14.00	2274	2188	45	9,33	4.67
11 •	23,850	72,690(1)	16.00	2197	2180	33	7.79	8,21
12	23,825	72.698.	16.00	2232	2203	33	7.79	8.21
13.	23.800~	72.700	16.00	2269	2217	50	9.85	6.15
14.	23.825	72,727	1800	2226	2245	55	10,33	7,67
15.	23,800/ 1:	72,727	18,00	2262	2257	52	10.05	7.95
16.	23.772	72.737	20,00	2301	2284	59	10,67	9.33
17.	24,266	72,183	14.50	1698	1257	45	9,33	5 .1 7
18.	24.1749	72,458	10,80	1 770	1698	35	8.09	2.71
19.	24.081	71,941	10,00	2027	986	26	6,61	3,39

CHAPTER - 3

MICRO DENSITOMETER METHOD

3.1 Micro Densitometer Approach:

3.1.1. General:

The micro-densitometer is an instrument which measures the light density at any point on the negative of a photo imagery. The Transmittivity, T, is defined as the ratio of the amount of light transmitted to the negative to the amount of light incident on the negative at any particular point on the negative. So 'T' can vary from '1' (transparent part of the negative) to '0' (Opaque part of the negative). Light Density, $D = \log_{10}(\frac{1}{T})$ --- (1).

Variations in surface features on land are reflected as variations in the shades or tones of a Black and White negative of an imagery. These variations in shades can be expressed in terms of variations in light density. Hence light density is an index of the properties of surface features mapped in an imagery. The microdensitometer gives a plot of transmittivity, T, versus distance along any straight line etc. on a photo - negative. The value of 'T' at any desired spot can be noted from this plot and the corresponding light density 'D' can be found out from equation (1).

3.1.2 Case Study:

Microdensitometer plots for this imagery were obtained from the Satellite Applications Centre, Ahmedabad (Gujrat). The photo negative was divided into six equi spacedlines such that three lines covered the Rajasthan region and the other three covered the Gujrat region.

Continuous Micro Densitometer plots for each of these lines were then obtained.

3.1.3 Calibration: Part:

A field study was undertaken to the Rajasthan and Gujrat regions and 50 wells were mapped and there water depths the latitudes, and longitudes of the nearest land mark on the map noted, as well as the orientation and the distance to the landmark from each well. The corrected latitudes and longitudes of each well were then determined from the following relation:

and the in to the Actual location from the assumed location

d = distance

Lat. corr. AØ"=± d Cost x 206265

 $R_{m} = 6378 \text{ Km.}$ (mean radius of earth)

These mapped ground water well points were then located on a ground water board map of the region (scale 1" = 4 mile). Those points which feel on any of the six lines (for microdensitometer study) or in the near vicinity, were chosen for calibration. The light density at these chosen points were ascertained from the micro-densitometer plots, for the Gujrat region (lines 4, 5, 6). Only the Gujrat region was considered as the CCT study had earlier indicated a lack of appreciable correlation in the Rajasthan region (see Figs. 2-5). The four plots were made on ordinary (see fig. No.11-14) graph paper by plotting light density versus depth. The first three plots were for light the light three plots were for light the light three plots were for l

Table - 8

Field data for Gujrat region
(Lines 4, 5, 6 Densitometer method)

	χ -	,				
Sl.	LINE	NO. 4	LINE 1	10. 5	↓ LINE	NO. 6
No.	Density	[Depth [(m.)	Density	Depth (m.)	Density	Depth (m.)
1.	0.796	5,60	0.745	5.90	0.921	5,40
2.	0,699	6.50	0,639	6.50	0,903	5,50
3.	0,620	8.00	0,620	7.20	0,757	5.60
4.	0.629	8.60	0.678	7.80	0.796	5.70
5 💂	0.602	9.30	0,602	8.90	0, 685	6,00
6.	0.620	10.00	0,585	9.20	0.678	6.30
7	0,620	10,00	0,602	9.50	0,638	7,10
8.	0.620	10.00	0.620	8.50	0,648	7,60
9.	0,629	9,80	0,620	7.60	0, 638	8,50
10.	0.620	9.40	0,658	6, 90		,
11.	0.638	9,00				
12.	0,678	7.70				

Assumed Latitude = d

Longitude = L

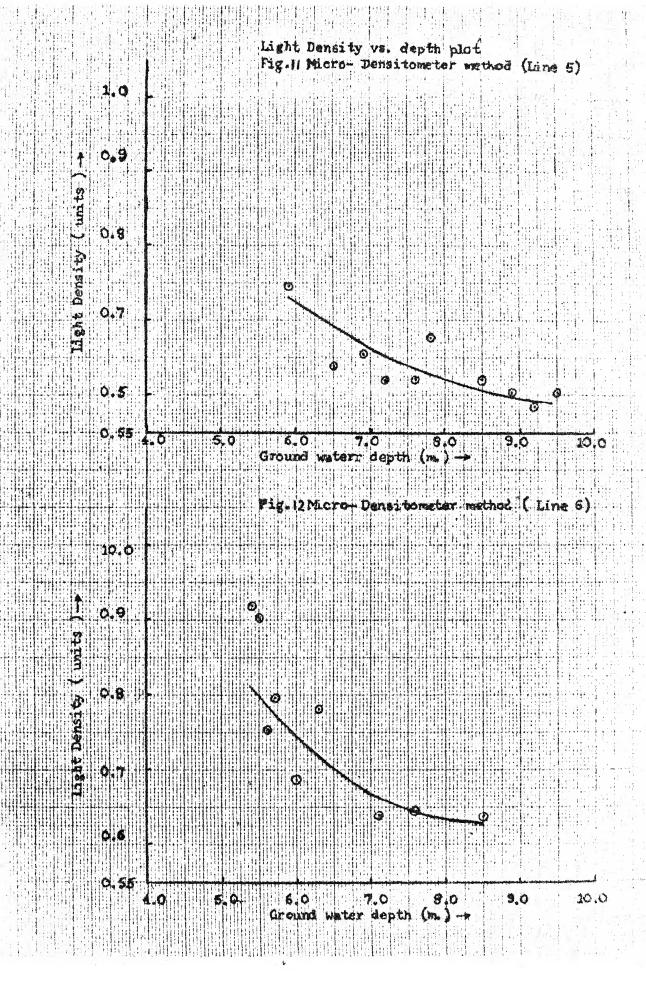
Longitude = L

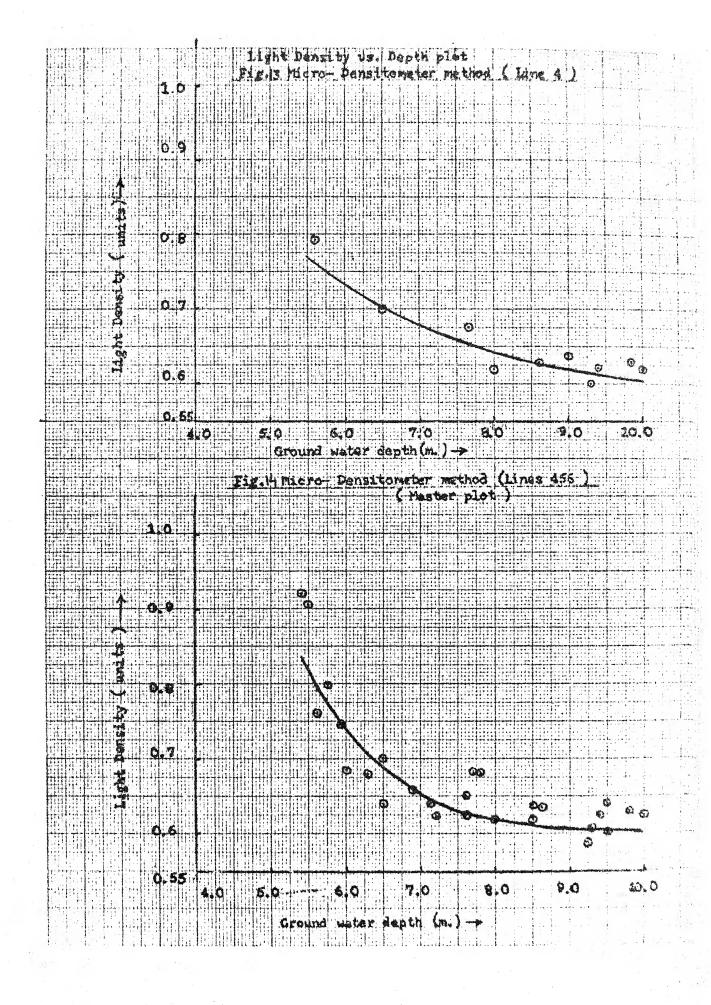
Location Corrected Lat, d' = d + 20"

(b, L)

(ovvected Lang, L' = L + 61"

Fig. 10





individually and the fourth plot was for the combined data of lines 4, 5, 6. Each of these four graphs established that at water depths deeper than 10 m there is so much 'scatter' on each graph that it can be concluded that there is no appreciable correlation beyond this depth, between the light density (satellite reading) and the water depth. Each graph showed an exponentially decaying curve of this mathematical form: $y = e^{bx} + c$. The equation of each curve was established by finding the values of 'b' and 'c' for each curve by the method of least squares curve fitting. Please refer to (Appendix - III)

The calibration curves for lines 4, 5, 6 and combined plot for lines 4, 5, 6 (master plot) are given in Figs. 11-14 and the curve equations and correlations coefficient (see appendix 3, are given in table 9.

Lines	Calibration B	Curve Equation B,C	Coefficient of Correlation
4. 5	- 0.04761652 - 0.04636945	- 0.02515450 - 0.09178096	- 0.88415922 - 0.78772666
6	- 0.10596548	+ 0.36929877	- 0.79501882
4, 5, 6 combined	- 0.05901143	+ 0.05047683	- 0.79258990

Table No. 9

Calibration Curve equation, $y = e^{-b x + c}$

3.1.4 Comments:

- i) The negative values of the coefficients of correlation, indicate that 'y' decreases with increase in 'x'.
- ii) Possible explanation for the light density to decrease with water depth is an exponentially decaying curve may be

as follows: water does not reflect radiation in Band - 7, hence a surface water body such as a lake, in Band - 7 would theoritically appear opaque on a positive film (i.e., on a 'positive' negative). Its transmittivity, T, would be '0' and, therefore, its density, D, would be infinite according to Eq.(1). As the ground water depth increases, the effect of soil moisture decreases and the same soil shows greater reflectance and, therefore, a greater transmittivity on a positive film, and hence a decreasing density. Hence a micro-densitometer plot of any line on a positive film of a particular area, in our case the Gujrat region, would show a high density value at very shallow ground water depth, which decreases exponentially with increase in the depth.

iii) The high value of 'r' indicates an excellent correlation between the light density and the water depth upto 10 metres. The plots for lines 4, 5 6 can be treated as auxilliary curves or verification curves for checking the water depths given by the master plot of lines 4, 5 6, though the latter by itself has a very high coefficient of correlation (-0.793). (The surface area covered by the master plot is approximately 185 Km x 90 Km. of Gujrat region). (See Figs. 11 to 14.)

3.2 Prediction Part (Micro-densitometer approach):

Only the Gujrat region the given imagery was considered for the prediction part because the given Rajasthan region indicated no appreciable correlation during the calibration phase (see 'CCT method' page 9).

In the CCT prediction process the given Gujrat region was subdivided into Regions (1) (2) & (3) (see page 19). In the densitometer method the Region (2) & Region (3) 'for lack of contour points in the given depth contour map) were combined as Region (2), & it was demonstrated that the satellite survey is not reliable for such regions (see'Results: densitometer method').

A number of points were taken on a 1" = 4 miles scale map of the Gujrat region supplied by the local ground water board authorities.

This map had ground water contours plotted on it by conventional methods.

Next each of these points were alloted to one of the three lines (lines 4,

5, 6) and the position of each point with respect to the corresponding line noted. Also the ground water depth at each point was recorded from this map, if necessary by interpolation.

Next the relevant line corresponding to each point on the map was considered and the corresponding light density value was noted from the microdensitometer plots. Then this density value was utilized to predict the corresponding ground water depth from the master ealibration curve as well as from the auxillary calibration curve for each line. The ground water depth(s) so obtained were then checked by comparison with the depths obtained from the conventional ground waster contour map (see Fig. 15). This completed the prediction of the ground water depth for one point on the map. This prediction process was repeated for the other points taken on the map for ground water prediction. The mean errors

Table - 10 ·

Field data and calculation for prediction:

(Gujrat region 1)

Sl.		l Latitude		l Density		Interated polated depth,D	Deviation x = P -D
	l (Degrees l	s)[(Degrees] [No.	Q Q Q	<pre>!depth !line 456 (m.),P</pre>	(m.)	(m.)
REG	ION 1						
1.	71.014	24.543	4	0,638	8.47	9,00	0.53
2.	71.840	24,520	4	0.593	9.71	9.27	0.44
3.	71,864	24.518	4	0.602	9,46	9.54	0.08
4.	71.884	24.514	4	0.595	9.65	9•82	0.17
5.	71.542	24.291	5	0.678	7.45	5.00	2,45
6.	71.579	24.284	5	04678	7.45	7,00	0,45
7.	71.626	24.277	5	0,638	8,47	7.00	1,47
8.	71,765	24.253	5	0.619	8.96	7.00	1.96
.9.	71,777	24.295	5	0,602	9.45	9,00	0,45
10.	71.412	24.044	6	0.638	8.46	7.00	1.46
11.	71,444	24.042	6	0,619	8.96	9,00	0.04
12.	71,535	24.026	6	0,638	8.46	9.00	0.45
13.	71.602	24.014	6	0;745	5.84	7.00	1,26
14.	71.693	24,000	6	0,745	5.84	7.00	1.26
15.	71.793	23.984	6	0,553	10.90	5.00	5,90
16.	71.884	23.969	6	0,638	8.46	7.00	1.46
17.	_71.479	24.040	6	0,699	6.92	8,00	1.08
18.	71.488	24.033	6	0,678	7.45	9.00	1.55
19.	71.576	24.019	6	0.854	3.53	8.00	4,47
20.	71,650	24.007	6	0 .67 8	7.45	6.00	1,45
21.	71.743	23.991	6	0,699	6.92	5.00	0, 08
22.	71.843	23,976	6	0.745	5.84	6.20	0.46
23.	71.600	24.014	6.	0.769	5.31	7.00	1,69
24.	71.800	23.980	6	0.699	6.92	5,00	1.92
				45 ***			

		Х					
Sl. No.	Longitu s e (Degrees)	Latitude (Degrees)		Density	Predicted depth line 456 (m.)P	Inter - polated depth,D (m.)	Deviation x=!P-D; (m.)
	-017		<u> </u>	^	A \\		
***************************************	ON 2						7 16
1.	72.000	24.512	4	0 .67 8	7.54	15 •00	7.46
2.	71.992	24.514	4	0.602	9,46	13,00	3.54
З.	71,983	24.517	4	0.600	9.51	11.00	1,49
4.	72.005	24.514	4	0.602	9,46	17.00	7.54
5.	72.011	24.512	4	0,638	8.47	19.00	10.53
6.	71.819	24.249	5	0.709	4.95	11,00	6,05
7.	71.851	24.239	5	0.658	7.96	13.00	5 _• 04
8.	71.893	24 225	5	0.678	7.45	15,00	7.55
9,	71.954	24.235	5	0,678	7.45	17.00	9.55
10.	71.954	24.228	5	0,638	8.47	19,00	10.53
111	•	24.221	5	0.638	8.47	19,00	10,53
12.	72.093	24.200	5	0.553	10.89	17.00	6.11
13.	72,219	24.179	5	0.769	5.29	17.00	11.71
14.	72.263	24.172	5	0,658	7.96	15.00	7.04
	72.302	24.165	5				, and the second
15.	72,321	24.184		0.658	7. 96	13,00	5.04
16.	72,339	24.160	5	0,619	8.97	11.00	2,03
17.	72.358	24.158	5	0 ,67 8	7.45	9,00	1,55
18.	72.388	24.149	5	0,585	9.94	9,00	0.94
19.	72,430	24.147	5	0.585	9.94	11.00	1.06
	P (g)						

and the standard errors were computed for Regions (1) and (2) to assess the accuracy of prediction in each region (see page 19).

Once it was established that this method provided a good accurate of prediction, the need of comparing with each predicted ground water depth with the conventional ground water contour map of an area is no longer necessary and the calibration prediction curves can be reliably used to predict (or update old conventional maps) shallow ground water depths in unmappedareas.

RESULTS AND DISCUSSION (HIGRO-DENSITYON FER APPROACH):

Region(1): (See table - 10)

Number of points N=24,

Standard error. $\sigma = 1.38 \text{ m}$

The 3 - O test shows that we have to reject any x>4.14 m hence we reject the two observations corresponding to serial nos 15 and 19 in table - 10. Therefore, we have :

N = 22, $\sigma = \pm 0.71$ m

We get a good prediction accuracy, therefore, in Region (1) as the standard error is 1.00 m. Region (1) is mostly Alluvial with ground water depths less than 10 m, where our satellite survey methods are most effective.

Re ion (2): (See table -11)

N = 19, 6= 3.56 m

The 3 - σ test shows that we have to reject any x > 10.68 m, hence one observed ion σ responding to serial no. 13 in table - 11 is discarded. Therefore, we get: $\Gamma = 16$, $\sigma = 3.39$ m.

The high value of the standard error indicate that the satellit survey methods are unreliable in Region(2), which is mostly hilly and with water depths greater than 10 m.

3.3 Sources of Error (CCT and Densitometer methods):

- 1. The normal water level in some of the wells in the Gujrat region was lowered by 0.5 m. 1.0 m. due to pumping operations, when the field data was collected. Hence the water levels in these wells represented a lesser values of the actual ground water depths.
- 2. Ideally, the sate: lite survey and the collection of field data (as well as preparing a conventional water depth contour map) should be done simultaneously. In practice, it is preferable to collect both types of data as close to each other as possible. Otherwise, the effect of seasonal variation of ground water depths comes into play, and the information recorded by a satellite and by a field observation, at any particular sport, will not indicate the same depth, but different depths due to seasonal variation. In this study, the satellite survey was done in June 1977, the field data was collected in December 1979 and February 1980, and the conventional water depth contour map was prepared in December, 1978. The correction due to seasonal variation could not be applied to the water depths due to lack of data. The low value of the standard error achieved (for ex.,+ 0.99 m. : CCT method), without correction due to the seasonal variation, indicates the high accuracy possible once the seasonal variation factor is taken into account.
- 3. To pin point the latitudes and longitudes, a 1" = 4 miles ground water contour map was used, for the CCT method.

 A higher accuracy could have been achieved with a larger scale map.

CHAPTER - 4

DISCUSSION. CONCLUSION AND RECOMMENDATIONS

4.1 DISCUSSION CONCLUSION AND RECOMMENDATIONS:

How well can ground water depth be estimated from satellite data?

For the given Rajasthan region, the application of the CCT method yielded no appreciable correlation between the ground water depth and the satellite data, using the central pixel method and the linear averaging method (3, 5 7 pixels array). This indicated that we can not reliably estimate water depths from satellite datain this region, which is mostly sandy and with water depths in excessof 10 m. The hilly terrain in the lower Rajasthan region, i.e. Mount Abu hills, were not included in this study because these prediction methods are not reliable in hilly areas.

Next we studied the Gujrat region given in the imagery with the help of the CCT method and the micro-densitometer method. Three different types of (for Regions (1), (2), (3), see page 19) were classified depending upon terrain (alluvial or hilly) and water depth (greater than or less than 10 m). The calibraion prediction yielded very promising prediction accuracy (atandard error=0.99 m). in case of CCT method andstandard error 0.71 m in case of densitometer method) (see pages 20, 40) in Region (1), where the terrain was mostly alluvial with water depths less than 10 m. It was not this type of region that we had the best results. On the other hand, as anticipated, both met ods yielded poor results (see pages 21, 40) for Regions (2) and (3), where the terrain was mostly hilly and water depths greater than 10 m.

A study of field data (see p. 29) in the Gujrat region, of depths greater than 10 m, established that the CCT method yields a poor accuracy (standard error = 2.63 m, see p.20).

Thus we can conclude that both the CCT method and the micro-densitometer met od are best applied to alluvial soil with shallow water depths not exceeding 10 m. A few points of interest are discussed here.

CCT Method Vis-a-vis the Micro-densitometer Method:

For the given Gujrat region, the correlation coefficients and the standard errors - 0.939 m and 0.99 m (Region (1)) iscase of the CCT met: od. and -0.793 m and 0.71 m. in case of micro-densitometer approach. It was found that ground water depths deeper than 10.0 m cannot be estimated because of a lack of correlation with the satellite data. Two points are to be noted. First isthat the above values of the correlation coefficients (0.8 - 0.9) indicate a strong correlation and the second point is that its value is higher for the CCT approach. (The correlation coefficients varies from-1) to +1 and the closer to unity it is, the stronger the correlation.) A correlation coefficient of 0-0.5 (test at 5% level of confidence) indicates no correlation to a negligibly small correlation while 0.7 - 1.0 indicate a good to an excellent correlation). Again the standard errors are very reasonable considering the sources of errors (see 'Sources of error'). The fact that depths could be predicted to a metre (CCT approach), points to the higher degree of precision possible with the minimization and removal of these sources of errors. The standard error is lesser for the CCT approach. The reason for this is that in CCT method, we have for more data

available (collected in lines and pixels) about every 80 m we have a reflectance reading - while in the microdensitometer method such precision was not available in the calibration part stage. The CCT approach is superior because it offers an enormously large wealth of data (1 line about a 185 km) contains 2656 pixels, each pixel about 80 m square). The next LANDSAT has a resolution of about 30 m and a French land use satellite is being planned for a resolution of 10 to 20 m. with these. With these satellites, the CCT method will offer an even more accurate picture of the ground water profile and an accuracy of within half a metre in areasconductive to satellite survey can be realized. This is not to belittle the microdensitometer method, which by itself is good enough to give good estimates but only that a higher precision can be obtained with the CCT method.

Central pixel method vis-a-vis the Grid averaging methods (CCT approach):

In the CCT method, in the study of Gujrat region, the correlation coefficient decreases from the central pixel method to 0.55 for the 7 pixels Grid pixel method and hence the degree of accuracy gets lesser. The reason for this trend is that we are assuming the ground water depth (which remains constant over a limited area only) to remains at a constant level over an increasing area as we move from the central pixel method to the 3 pixels, 5 pixels and 7 pixels Grid averaging. The precision therefore decreases. In the 7 pixels land case, the correlation coefficient is only 0.55, indicating a very poor correlation. This means that the ground water depths should not be averaged beyond a ground surface distance of about 400 m allowed

the 5 pixels grid method (each pixel > 80 m, length of a 5 pixels grid \pm 5x80 = 400m). Otherwise there will result a negligibly small correlation or no correlation as in case of a 7 pixel grid averaging of the Gujrat area. Also it means that the central pixel method is most accurate in the CCT approach. The restriction of 400 m. on the grid averaging also means that a remote sensing satellite should have atleast a pixel resolution of 400 m to map shallow ground water depths. The first Indian Remote Sensing Satellite, Bhaskara - 1, has a pixel hence lacks the precision required. It is much larger hoped that with the Indian Space Research program maintaining its ambitious program, ISRO will in a not too distant future put a Bhaskara series satellite of pixel resolution well below 100 m. When this happens indigeneous satellite survey based on the CCT and microdensitometer methods explained in this study will be realized. Till we achieve the required pixel resolution in our very own Bhaskara etc., ISRO can procure LANDSAT data to develop a satellite ground water depth contour map of most parts of India.

Sotellite Survey vis-a-vis Aircraft Survey:

Aircraft Survey: Versus Satellite Survey

The objective of ground water depth prediction can be achieved using both an aircraft and the satellite with MSS equipment. But the advantage of a satellite is the larger area coverage (185 Km x 185 Km: LANDSAT) compared to an aircraft. A satellite survey more suitable in a ground water depth estimation in large areas, which would require a large number of flights (and camera exposures) by aircraft. Aircraft survey is suited for covering small local areas or for those projects where steroscopic photographs are needed.

It is obvious that to prepare ground water maps of large areas, the superior option is a salellite borne survey.

Summing it up, this study has proved that it is possible to predict shallow ground water depths from remote sensing satellites with a good accuracy. Ground water depths upto 10 m. depth can be estimated while rocky areas or built up areas or areas under dense forest cover etc. cannot be covered by a satellite surveys. The given Gujrat region was probed with two techniques, one using CCTs and the other a micro-densitometer. Upto 10 m. depths could be estimated with a high correlation and very reasonable standared errors especially when quite a few sources of error existed. Undoubtably the standard error of ± 0.99 m.(CCT approach) can be refined considerably with the removal of these sources of error.

The highly encouraging results obtained in this project highlight the tremendous scope of predicting shallow ground water depths by satellite surveys.

```
THIS PROGRAM CONVERTS THE LATITUDE AND LONGITUDE OF A PUINT ON THE GROUND TO ITS, CORRESPONDING LINE-NUMBER AND PIXEL NUMBER ON THE IMAGERY AFTER CORRECTING FOR THE CURVATURE OF THE EARTH.
                            IMPLICIT DOUBLE PRECISION (A-R,0-Z)
DIMENSION PL(200), PPHI(200), XP(200), YP(200), U(200), V(200)
9, OPHI(200)
OPEN(UNIT=7, DEVICE='DSK')
READ(7,*)N
READ(7,*)N
READ(7,*)N
FORMAT(14)
FORMAT(2F11.6)
DO 905 I=1, N
OPHI(1)=PPHI(1)*3.141593/180.
CALL VAL(RPHI, R, OPHI, OL, RM, RN)
DL=PL(1)-UL
DPHI=OPHI(1)-OPHI
 100
 200
                             DL=PL(I)-UL

DPHI=OPHI(I)-OPHI

C=DL*(SIN(OFHI))

SM=RM*(UPHI(I)-UPHI)

OM11=SM**3

DM12=6*(R**2)

DM12=6*(R**2)

DM12=1-5M**4

DM212=SIN(OPMI)/CGS(OPFI)

OM212=SIN(OPMI)/CGS(OPFI)

OM22=DM211*DM1-DM2

RPDM=SM+DM1-DM2

RPDM=RPHI-DM2
                             LATITUDE=
1000
                                                                                                                                                                                ,F6.0)
  905
                              SUBROUTINE VAL(RPHI, P.OPHI, OL, RM, RN)
F=6.37730124
G=10**6
A=F*G
E=1/300.8
                              E=1/300.8

OPHI=23.*3.141593/180.

H=E**2

S=SIN(OPHI)

S2=S**2

B=SORT(1-H*S2)

RH=A/B

D=B**3

UN=A*(1.-H)

RM=UN/D

R=SORT(RN*RM)

RPHI=R*(COS(OPHI)/SIN(OPHI))

OL=71.

RETURN

END
```

```
PROGRAM TO READ THE REFLECTANCE VALUES IN BINARY FORM FROM THE TAPE AND THEN TO CONVERT THEM TO THEIR DECIMAL VALUES.

INTEGER BT1,BT2,BT3,BT4,BT5
DIMENSION 1A(840),LAST(30),VL(30)
OPEN(UNIT=20,DEVICE='MTAO',MDDE='DUMP'
9,RECORD SIZE=840,DENSITY='800')
OPEN(UNIT=21,FILE='FGR21.DAT')
DATA 1B/5H
DATA LAST/30*999/
ACCEPT*,MMM,M
FORMAT(5X,' THE LINE NUMBER OF THIS RECORD IS',I5)
PRINT 61,MMM
NI=M+12
KAT=0
61
                         NI=M+12
KAT=0
K=0
NREC=0
D055N=1,840
IA(N)=IB
READ(20)IA
NO=-3
K-V-1
99
55
                         K=K+1
IF(K.GT.1)GO YO 9U
I=0
                       221
25
500
112
88
111
888
 90
```

```
PROGRAM TO FIND CURVE EQUATION DIMENSION X(80), Y(80), A(80)
                                                               NUNFO
                                                               4=1
                                                               N=75
                                                            N=75
READ(21,*)(X(1),T=1,N),(Y(J),J=1,N)
PRINT 44,N
FORMAT(15X,'INPUT DATA FOR CALIBRAT
75X,'IN=',1,2)
PRINT 770,X(1),Y(1)
PORMAT(10X,'X(1)=',F15.8,5X,'Y(1)='
CONTINUE
N=N-44
CONTINUE
C=0.0
ANF=N-M+1
DO 1 1=M,N
                                                                                                                                                    INPUT DATA FOR CALIBRATION-PREDICTION
44
                                                                                                                                                                                                            ,F15.8,5X, Y(I)=",F15.8)
42
5
                                                              DO 1 1=M, N
Z=Y(I)
A(I)=ALOG(Z)
C=C+A(I)
COSTINUE
                                                              J=M,N
                                                                D=D+DD
                                                              B=B+W
AA=W*W
TA=TA+
                                                                      A=W*W
A=TA+AA
OMTINUE
P(NUN.GE.4)GOTO 111
UN=NUN+3
F(MUN.LE.3)MUN=457
PRINT 400,MUN
ORBAT(///, 10X, 'MICRO-DENSITOMETER METHOD', 10X, 'JUNETINE NUMBER(S)=', 12)
TO TO 10
TO TO 10
2
                                                           | IF (MUN, LE. 3) MUN=457
| PRINT 400, MUN | FORMAT(///, 10x, MTCFO-DENSITOMETER |
| 4*LINE NUMBER(S)=', 12) |
| GD TO 10 |
| CONTINUE | NIM=NUN-5 |
| LIN=NIM+2 | PRINT 401 | LIM |
| FORMAT(///, 10x, MAGNETIC TAPE METHO |
| FORMAT(///, 10x, MAGNETIC TAPE METHO |
| FORMAT(///, 10x, MAGNETIC TAPE METHO |
| CONTINUE | DESCRIPTION |
| BB=(B*C-ANF*D)/(B**2-TA*ANF) |
| CC=(B*D-TA*C)/(B**2-TA*ANF) |
| PRINT100.BB,CC.NUM |
| FORMAT(//, 5x, BB=', F15.8, 2x, CC=', F10 |
| CC=(B*D-TA*C)/BB |
| CC=(
  400
 111
 401
                                                                                                                                                                                                                                                                       TAPE METHOD', 10X.
                                                                                                                                                                                                                                                                                                                                                                                                                   'NUMBER OF
 10
                                                                                                                                                                                                                       ,F15.8,2X, 'CC=',F15.8,2X, 'NUN=',I2)
  100
 200
   1
  110
```

```
PRINT 700, AVX, AVY, NUN
FORMAT(//,5X, 'AVX=',F15.8,2X,'AVY=',F15.8,2X,'NUN=',I2)
2SUM=0.0
YSUM=0.0
DO 50 I=M, N
0=X(I)-AVX
R=0**2
S=A(I)-AVY
V=S**2
A0=0**
PRINT 500, X(I), Y(I), R,S,AQ
FORMAT(/,5X,'X(I)=',F10.4,2X,'Y(I)=',F6.3,2X,'RX=',F15.8,'XSUM=XSUM+R
YSUM+YSUM+V
ZSUM=ZSUM+AQ
CONTINUE
IF(J,LT.30)J=J-1
AX=XSUM/J
AX=YSUM/J
AX=YSUM/J
XSTD=SORT(AX)
YYSTD=SORT(AX)

700
    500
    50
    600
                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                                 'XSUM=',
                                                                                                                                                                                                                                                                                                                                            12 TIN SE

12 TIN SE

12 TIN SE

13 TIN SE

14 TIN SE

15 ON 13 SE

16 ON 13 SE

17 ON 13 SE

18 ON 15 SE

18
    7
    8
    11
    12
    13
         14
```

88=	-0.05	901143 CC	=	0.05047683	NUN= 0	
	X(I) =	5.6000	XX=	4.7217	XDIF=	0.8783
	Y(I)=	0.7950	YY=	0.7558	YDIF=	0.0402
	X(I)=	6.5000	xx=	6.9238	XDIF=	-0.4238
	Y(I)=	0.6990	YY=	0.7167	YDIF=	-0.0177
	X(1)=	8.0000	XX=	8,9561	XDIF=	-0.9561
	Y(I)=	0.6200	Y Y =	0.6560	YDIF=	-0.0360
	X(I)=	8,6000	XX=	8.7119	XDIF=	-0.1119
	Y(I)=	0.6290	YY=	0.6332	YDIF=	-0.0042
	x(I)=	9.3000	XX =	9.4554	XDIF=	-0.1554
	Y(I)=	0.6020	Y Y =	0.5075	YDIF=	-0.0055
	X(I) =	10.0000	XX=	8.9561	XUIF=	1.0439
	Y(I)=	0.6200	YY=	0.5830	YDIF=	0.0370
	X(I)=	10.0000	XX=	8.9561	XDIF=	1.0439
	Y(I)=	0.6200	Y Y =	0.5830	YDIF=	0.0370
	X(I) =	10.0000	XX=	6.9561	XDIF=	1.0439
	Y(I)=	0.6200	X X =	0.5830	YOIF=	0.0370
	X(I)=	9.8000	XX=	8.7119	XDIF=	1.0881
	Y(I)=	0.6290	YY=	0.5899	YDIF=	0.0391
Alpho Alpho	X(1)=	9,4000	xx=	8.9561	XDIF=	0.4439
	Y(I)=	0.6200	YY=	0.6040	YDIF=	0.0160
	X(1)=	9.0000	xx=	8.4711	XDIF=	0.5289
	Y(I)=	0.6380	YY=	0.6184	YDIF=	0.0196
	X(I)=	7.7000	XX=	7.4407	XDIF=	0.2593
	Y(I)=	0.6780	YY =	0.6677	XDIF=	0.0103
2 th	X(I)=	5.9000	Xx=	5,8437	XDIF=	0.0563
14.21	Y(I)=	0.7450	YY=	0.7425	YDIF=	0.0025
	x(I)=	6.5000	xx=	8,4446	XDIF=	-1.9446
	Y(I)=	0.6390	YY=	0,7167	YDIF=	-0.0777
	X(I)=	7.2000	XX=	8.9561	XDIF=	-1-7561
	Y(I)=	0.6200	YY=	0.6877	YDIF=	-0.0677
	X(I)=	7.8000	XX=	7.4407	XDIF=	0.3593
, A	Y(1)=	0.6780	YY=	0.6638	YDIF=	0.0142
	*(1)=	8.4000	XX=	9.4554	XDIF=	+0.5554
	Y(1)=		the second second		YOTF=	-0.0201
	X(I)=		XX=		XDIF=	-0.7408

```
-0.0261
Y(I) =
         0.5850
                   YY=
                          0.6111
                                   YDIF=
                                             0.0446
                  XX=
                          9.4554
                                   XDIF=
         9.5000
X(I) =
                                   YDIF=
                                            0.0016
                          0.6004
         0.6020
                   YY=
Y(I)=
                                            -0.4561
                                   XDIF=
X(T) =
         8.5000
                   XX=
                          8.9561
                                            -0.0169
                                   YDIF=
Y(I)=
         0.6200
                   XX=
                          0.6369
                                            -1.3561
                                   XDIF=
                          8.9561
X(I) =
         7.6000
                   XXX
                                            -0.0517
                                   YDIF=
Y(T) =
         0.6200
                   YY=
                          0.6717
                                            -1.0481
         6.9000
                          7.9481
                                   XDIF=
                   XX=
X(I) =
                                            -0.0420
                          0.7000
                                   YDIF
Y(I)=
         0.6580
                   YY=
                          2.2499
                                   XDTF=
                                             3.1501
         5.4000
                   XX=
X(I) =
                                             0.1562
                          0.7648
                                  YDIF=
         0.9210
                   YY=
Y(I) =
                                             2.9156
                          2.5844
                                   XDIF=
         5.5000
                   XX=
X(I) =
                                             0.1427
                          0.7603
                                   YDIF=
Y(I)=
         0.9030
                  YY =
                                             0.0270
                          5.5730
                                   XDIF=
X(I) =
         5.6000
                   XX=
                                             0.0012
         0.7570
                          0.7558
                                   YDIF=
Y(I) =
                   YY=
                          4.7217
                                    XDIF=
                                             0.9783
         5.7000
X(I) =
                   XX=
                                            0.0447
Y(I) =
         0.7960
                   YY=
                          0.7513
                                    YDIF=
                                    XDIF= -1.2666
                          7.2666
X(I)=
         6.0000
                   XXX
                                    YDIF= -0.0532
                          0.7382
         0.6850
                   YY=
Y(I)=
                          7.4407
                                    XDIF=
                                            -1.1407
          6.3000
                   XX=
X(I) =
                                    YDIF= -0.0472
          0.6780
                   YY=
                          0.7252
X(I)=
                                            -1.3711
                          8.4711
                                    XDIF=
          7.1000
X(I)=
                   XX=
                          0.6918
                                            -0.0538
                                   YDIF=
          0.6380
                   YY=
Y(I)=
                          8.2076
                                    XDIF=
                                            -0.6076
          7.5000
                  XX=
X(I)=
                                            -0.0237
                                    YDIF#/
                   YY=
                           0.6717
Y(I)=
          0.6480
                                             0.0289
                         8.4711
                                    XDIF=
X(X) =
          8.5000
                   XX=
                                              0.0011
                                    YDIF=
                           0.6369
Y(I)=
          0.6380
                   YY=
```

XSTAN. DEV.= 1.521 YSTAN. DEV.= 0.113

MICRO-DENSITOMETER METHOD

LINE NUMBER(S)=

BB= -0.04	761652 ČC	:	0.02515450	NUN= 1	4
X(T)=	5,6000	XX=	4.2633	XOJF=	1,3367
¥(1)=	0.7960	YY=	0.7469	YDIF=	0.0491
X(1)=	6.5000	XX=	6.9923	XDIF=	-0.4923
Y(1)=	-0.6990	YY=	0.7156	YDIF=	-0.0166
×(1)=	8.0000	2X=	9.8110	XDIF=	-1.5110

```
-0.0463
                          0.0663
                                    YDIF=
=(I)Y
         0.6200
                   YYF
                                    XDIF=
                                            -0.6083
         8.6000
                   XX=
                          9.2083
X(I) =
                                            -0.0185
         0.6290
                          0.6475
                                    YDIF=
Y(I)=
                   YY#
                                             -0.8297
         9.3000
                   XX
                         10.1297
                                    XDIF=
X(I)=
                                    YDIF=
                                             -0.0243
         0.6020
                   YY=
                          0.6263
Y(T)=
                                             0.4890
        10.0000
                   XX=
                          9.5110
                                    XDIF=
X(I)=
                          0.6057
                                    YDIF=
                                             0.0143
         0.6200
                   YY=
Y(T)=
                                              0.4890
                          9,5110
                                    XDIF=
        10.0000
X(I) =
                   XX =
                                    YD IF=
                                             0.0143
         0.6200
                   YY=
                          0.6057
Y(I)=
                                              0.4890
        10.0090
                          9.5110
                                    XDIF=
                   XX=
X(I) =
                                              0.0143
                                    YOIF=
Y(I) =
          0.6200
                   YY=
                           0.6057
                                              0.5917
                           9.2063
                                    XDIF=
          9.8700
                  XX=
X(I) =
                          0.6115
                                    YOIF=
                                              0.0175
          0.6290
                   YYZ
Y(I) =
                                             -0.1110
                           9.5110
                                    XDIF=
          9.4000
                   XX=
X(I) =
                                             -0.0033
                           0.6233
                                    YDIF=
Y(I)=
          0.6200
                   YY=
                           8.9100
                                    XOTF'=
                                              0.0900
          9.0000
                   XX=
X(I) =
                                              0.0027
                                    YDIF =
          0,6380
                   YY=
                           0.6353
X(I)=
                                              0.0671
                           7.6329
                                    XDIF=
          7.7000
                   XX=
X(I) =
                                    YDIF=
                                              0.0022
                           0.6758
                  YY=
Y(1)=
          0.6780
```

CORRELATION COEFFICIENT TO 88415922

#ICRO-DENSITOMETER METHOD LINE NUMBER(S) = -0.04636945 CC= -0.09178096 NUN= 2

+0.04	536945 CC		-0.031/8030	MUM=	
X(I)=	5.9000	XX=	4,3690	XDIF=	1.5310
Y(3)=	0.7450	YY=	0,6939	YDIF=	0.0511
x(I)=	6.5000	XX=	7.6790	XDIF=	-1.1790
Y(1)=	0.6390	γγ=	0.6749	YDIF=	-0.0359
X(I)=	7.2000	XX=	8,3299	XDIF=	-1.1299
Y(I)=	0.6200	XX=	0,6534	YOLF=	-0.0334
X(I)=	7.8000	XX#	6,4013	XDIF=	1.3987
Y(I)=	0.6780	YY=	0,6354	YDIF=	0.0426
X(T)=	8.9000	XX=	9.9653	XDIF=	-0.0653
X(I)=	0.6020	YY=	0.6038	YDIF=	-0.0018
Y(I)=	9,2000	XX=	9,5831	XOIF=	÷0.3831
Y(I)=	0.5050	YY=	0.5955	YOIF=	+0.0105

```
X(I) = 9.5000
                    8,9653 XDIF= 0.5347
              XX=
Y(I)=
      0,6020 YY
                   0,5873
                           YDIF=
                                   0.0147
X(I) = 8.5000 XX =
                   . 9.3299 XDIF= 0.1701
Y(I) =
      0.6200 YY=
                   0.6151 YDTF= 0.0049
X(I) = 7.6000 XX =
                           XDIF= -0.7299
                   8,3299
      0.6200 YY=
                   0.6413 YDIF= -0.0213
Y(1)=
X(I)=
      5.9000
             XX=
                    7.0471 XOIF= -0.1471
                   0.6625 YDIF=
Y(I) = 0.6580 YY=
                                  -0.0045
```

CORRELATION COEFFICIENT= TO 178772666

MICRO-DESSITOMETER METHOD

LINE NUMBER(S)=

BB=	-0.1	.0596548 CC	=	0.36929877	NUN=	3
	X(1)=	5.4000	XX=	4.2617	XDIF=	1.1383
	Y(I)=	0.9210	XX=	0.8163	YDIF=	0.1047
	×(1)=	5.5000	XX=	4.4480	XDIF=	1.0520
	Y(I)=	0,9030	YY=	0.6077	YDIF=	0.0953
	x(I)=	5,6000	XX=	6.1123	XDIF=	-0.5123
	YCI)=	0,7570	X X =	0.7992	YDIF=	-0.0422
	X(I)=	5.7000	XX=	5.6382	XDIF=	0.0618
,	F(T)F	0.7960	YY=	0.7908	YDIF=	0.0052
	X(I)=	6.0000	XX=	7.0555	XDIF=	-1.0555
	Y(1)=	0.6850	YY=	0.7661	YDIF=	-0.0811
	X(T)=	6.3000	XX#	7,1524	XDIF=	-0.8524
	Y(I)#	0.6780	YY#	0.7421	YDIF=	-0.0641
	X(I)=	7.1000	xx≠	7.7262	XDIF#	-0.6262
	Y(I)=	0.6380	YY=	0.6818	YDIF=	-0.0438
3	x(1)=	7,6000	XXŦ	7.5795	XDIF=	0.0205
	Y(I)=	0.5480	YY=	0,6466	YDIF=	0.0014
	X(I)=	8.5000	XX=	7.7262	XDIF=	0.7738
	Y(I)=	0.6380	XX=	0.5879	YDIF=	0.0502
	- officially four		The Market of the second			The second second

CORRELATION COEFFICIENT +0.79501882

BB=	0.20	175720	CC=	1,92411720	NUN=	4
	X(1)=	9.9000	xx=	10.5886	XDJF=	-0.6886
	Y(I)=.	58,0000	YY=	50.4768	YDTF=	7.5232
	X(I)=	4.6000	XX=	5.0572	XDIF=	-0.2572
	Y(I)=	19.0000	YY=	18.0393	YDIF=	0.9607
	X(1)#	6.6000	XX=	6.7989	XDIF=	0.0011
	(I)=	27.0000	YY=	27,0062	YDIF=	-0.0062
	X(I)=	5.0000	XX=	5.0572	XDIF=	-0.0572
	Y(I)=	19.0000	YY=	18.7821	YDIF=	0.2179
	X(I) =	5.8000	X X =	5.5532	XDTF=	0.2468
	Y(I)=	21.0000	YY=	22.0720	YDIF=	-1.0720
	X(I)=	7.7000	XX =	6.2151	XDIF=	1.4849
	Y(I)=-	24.0000	YYE	32.3835	YDIF=	-8.3835
	X(I) =	5.4000	XX=	6.2151	XDJF=	-0,8151
	Y(I)=	24,0006	Y Y =	20.3607	YDIF=	3.6393
	x(1)=	10.0000	XX=	10.6733	XDIF=	-0.6733
	Y(I)=	59.0000	Y Y =	51.5055	YDIF=	7,4945
	*(I)=	8.7000) XX=	7.9414	XDIF=	0.7586
	Y(I)=	34.0000) , yy=	39,6228	YOIF	-5.6228
	Carlo Maria					The state of the s

CORRELATION CHEFFICIENT YSTAN. DEV. = 0.440

MAGNETIC TAPE METHOD

NUMBER OF PIXEL(S)= 2

				1)			- 42.1	100					独身 文	a Albert	1011 TO			100
68= ,		0.1	60	10	59	9	CC=		2	.41	2945	30	NU	N=	5			
	XII)=		9	. 9	000		XX=		10.	3968		XDI	F'=		-0.	49	68
	Y(I)=		59	.0	000		YYF		54.	4888		YDI	F=		4.	51	12
	XCI)=		4	. 8	000		XX=		4.	2353		XDI	F=		0.	564	47
	YCI) =	6/	22	.0	000		YY=		24.	0817		YDI	F=		-2.	08	17
	X(I)=	13	6	. 8	000	11.	XX=		5.	2787		XDI	F'=		1.	52	13
	XCI) =		26	.0	000		Y.Y.=		33.	170	-	YDI	F=		-7,	17	07
	X(I)=		5	. 0	000	+30	XX#		5.	514		XDI	F'=		-0.	51	44
	YCI)=		27	.11	000		YY=		24,	865	3	YDI	F=		2.	. 13	47
	X(I)=		5		000		XX÷		4.	7781		XDI	F=		1.	02	12
	YLI)=		24	.0	000		YYF		28,	263	2	YDI	F=		-4.	26	32
	XCI)=		7	. 7	000		XX=		7,	611		XDI	r=		-0.	.11	12
					de la companya de la	n, P												

Y(I)=	39,0000	YYP	38,3119	YDIF=	0.6881
X(I)=	5.4000	XX=	7.4824	XDIF=	-2.0824
Y(I)=	37.0000	YY=	26,5099	YDTF=	10.4901
X(I)=	10.0000	XX=	10.1814	XDIF=	-0.1814
Y(I)=	57.0000	YY=	55.3682	YDIF=	1.6318
K(T)#	8.7000	XX=	8.4210	XDIF=	0.2790
Y(I)=	43,0000	YY=	44.9642	YDIF=	-1.9642

XSTAN. DEV.= 2.047 YSTAN. DEV.= 0.368
CORRELATION COEFFICIENT= 0.89163342

MAGNETIC TAPE METHOD NUMBER OF PIXEL(S) = 3

88=	0.0	7443128	CC=	3.08453320	NUNF	6
	x(T) =	9,9000	XX=	10.2862	XDIF=	-0.3862
	Y(I)=	47.0000	YY=	45.0682	YOIF=	1.3318
	X(I)=	4.8000) XX=	4,2545	XDIF=	0.5455
	Y(I)=	30.0000	YY=	31.2432	YDJF=	-1.2432
	X(I)=	6.8000	XX=	5.1215	XDIF=	1.6785
	Y(I)=	32,0000	YY=	36.2582	YOIF=	-4.2582
	X(I)=	5.0000	=XX	4.6950	XDIF=	0.3050
	=(I)Y	31.000	=XY	31.7118	YDIF=	-0.7118
	X(1)=	5.800	=xx=	4,6950	XDIF=	1.1050
	Y(I)=	31.000	= XX	33,6575	YDIF=	-2.6575
	x(I)=	7.700	=XX=	8,4513	XDIF=	-0.9513
	Y(I)=	41.000	= YY	38,7703	YOIF=	2.2297
	X(I)#	5,400	xx=	0,1195	XDIF=	-2.7195
	A low the state of the same	K. T.			ALLEN CO.	

Y(I)=	40.0000	YY=	32,6702	YDIF=	7.3298
X(I)=	10.0000	XX=	9,7020	XDIF=	0.2980
Y(I)=	45.0000	- YY=	46,0094	YDTF=	-1.0094
X(I) =	8,7000	•XX=	8.7750	XOTF=	-0.0750
Y(I)=	42.0000	YY=	41.7661	YDIF=	0.2339

CORREGATION COEFFICIENT= VSTAN DEV = 0.179

MAGNETIC TAPE METHOD NUMBER OF PIXEL(S) = 4

BB=	0.0	3375640	CC=	3.46698590	NIIN	7	* * * 1 *
	X(X) =	9.9000	XX=	8.6566	XDIF=		1.2434
	X(I) =	43.0000	YY=	44.8433	YDIF=		-1.8433
	X(I) =	4.8000	XX=	3,3929	XDIF=		1.4071
	Y(1)=	36.0000	YY=	37', 7512	YDIF=		-1.7512
	X(I) =	6.8000	XX=	5.7641	XDIF=		1.0359
	Y(I)=	39,0000	YY=	40.3879	YDIF=		-1.3879
	X(I)=	5,0000	XX=	2.5584	XDIF=		2.4416
	Y(I)=	35.0000	YY=	38,0069	YOIF=		-3.0069
	X(I)=	5,8000	XX=	2.5584	XDIF=		3,2416
î.	Y(T)=	35.0000	YY=	39.0473	YDIF=		-4.0473
	X(I)=	7.7000	XX=	10,0033	XDIF=	1 S	-2.3033
, W	Y(I)=	45.0000	YY=	41,6337	Yote=		3.3663.
1	X(I)=	5.4000	XX=	11.2915	XDIF=	T.	-5.8915
	¥(1)=	47.0000	YY=	38,5236	YDIF=		8,4764
	X(I)=	10,0000	XX=	7.9595	XDIF=		2.0405
O Allegan	Y(I)=	42.0000	YY=	44.9949	YDIF=		-2,9949
	X(I)=	8.7000	XX=	11,9152	XDIF=		-3.2152
	4(1)=	48,0000	YY=	43.0631	YDIF=		4.9369
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APPENDIX-V
                                                      APPENDIX-V
DATA:RAJASTHAN REGION(CCT METHOD)
LONGITUDE LATITUDE LINE NO. PIX
(DEGREES) (DEGREES)
71.810000 25.350000 179 24
72.252778 25.244444 235 92
72.516666 25.202778 236 13
72.466666 25.202778 248 12
72.425000 25.200000 261 11
72.388889 25.183333 271 12
                 FIELD
  S.NO.
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                                                        71.650000
72.020000
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 41.
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 42.
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  43.
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KEFERENCES

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